



# **Creep of frozen soils**

**Prof. Jilin Qi** 

## **College of Civil and Transportation Engineering, Beijing University of Civil Engineering and Architecture**





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### AGENDA

- □ Factors influencing creep of frozen soils
- □ Creep of Warm Frozen Soils: Field Observations
- □ State of The Art: Creep Models for Frozen Soils
- Our attempts on constitutive modeling
- Concluding Remarks





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# **CREEP CURVES FOR FROZEN SOILS**



TYPICAL creep curves for frozen soil (Ting 1983)

Frozen soil displays features very similar to those of unfrozen soil





# **STRESS DEPENDENCE**





Uniaxial compression Frozen Lanzhou sand From Wu and Ma. 1994 Triaxial compression Frozen ISO Standard sand

### Both very much stress dependent.





## **S**TRAIN RATE DEPENDENCE









Temperature plays a role similar to load Minimal strain rate and time to failure all dependent on temperature





# **DEPENDENT ON TOTAL WATER CONTENT\***



**<u>1.</u>** -0.5 °C, **1.4 MPa; 2.** -1.0 °C, **2.2 MPa; <u>3.</u>** -2.0 °C, **3.4 MPa** From Wu and Ma. 1994

Generally, different water content range, different minimal strain rate development for frozen silty clay (samples from Lanzhou, China)

\* Total water content: ice is counted as water together with unfrozen water





# **SOLUTE DEPENDENCE**



The higher solute content, the larger axial strain under a certain stress level.





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# **DEFINE WARM FROZEN SOIL**

#### **State-of-the-art:**

- Shields et al. (1985): -2.5 and -3.0 °C.
- Foster et al. (1991): -1 °C
- Tsytovich(1975): different temperature boundaries for different soils

### **Temperature boundary:**

- Detecting unfrozen water content is rather difficult (NMR)
- Tend to use mechanical properties

Our Work :

### Constant load and stepped temperature

- Material: silty clay (CL) Qinghai-Tibet Plateau
- Different water content: 40%, 80%, 120%
- Constant Loads: 0.1, 0.2 and 0.3 MPa, respectively
- Temperature steps: -1.5, -1.0, -0.6, -0.5 and -0.3 °C





### **Strain rate vs. Temperature**



like we applied higher loads in oedometer test





# **Compressive Coefficient vs. Temperature**



-1 °C is defined as the temperature boundary for warm frozen soil for the silty clay we frequently encounter on the plateau.





# Settlement of road embankment







# FIELD OBSERVATION QINGHAI-TIBET RAILWAY (DR. JIANMING ZHANG)



# FIELD OBSERVATION QINGHAI-TIBET HIGHWAY







### Along the highway: 300 km, 10 observation sites, 4500 m a.s.l.









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# **CREEP MODELS FOR FROZEN SOILS**

Model classification based on its scale of representation:

# • Microscopic view

- > The theory of rate process
- > Damage creep model

## • Phenomenological view

- Empirical model
- Elementary element based creep model





# CREEP MODELS FROM MICROSCOPIC VIEW\_1

> The theory of rate process: Considers deformation as a thermal activation process



Related studies: Andersland (1967), Assur (1980)

#### A reasonable physical description

**Difficulty:** Parameters obtained qualitatively by current testing technology





# CREEP MODELS FROM MICROSCOPIC VIEW\_2

#### The damage creep model: damage or recovery of soil structure



From Miao et al. 1995

Following thermodynamics; Unique internal variable in frozen soil: ice content *Difficulty*: Calibration of Parameters for thermodynamics and damage mechanics





# CREEP MODELS FROM PHENOMENOLOGICAL VIEW\_1

Empirical methods: a mathematical description of creep curve

#### Basic form:

Classified by creep stages:



- I. Primary creep model (Vyalov, 1966)
- II. Secondary creep model (Ladanyi,1972)
- III. Tertiary creep model (?)

Simple structure; conveniently applied in simple engineering analysis (first estimation)

1) Poor versatility; 2) do not reflect internal mechanism





# CREEP MODELS FROM PHENOMENOLOGICAL VIEW\_3

#### Elementary creep model: combination of mechanical elements



Reasonable mechanical basis; simple structure; convenient in engineering design

Reasonable parameters decide the precisions.





# SUMMARIZATION

- Some are too complicated in form, too many parameters, even impossible to be obtained from conventional tests
- Some are lacking mechanism, just mathematical description
- Some are difficult to accommodate different thermal or load conditions

We need something new.





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# BORROW THEORIES FROM UNFROZEN SOILS

- Ladanyi (1999): creep of frozen soils is not very much different from that of unfrozen soils
- Warm frozen soil is between frozen and unfrozen soils, most likely closer to unfrozen soils





### –Creep of unfrozen soils based on $p_c$



(After Bjerrum, 1973)





# Questions

### Is there an index in frozen soils similar to $p_c$ ?

If so

What is its relationship with creep?

How is it possible to apply creep model of unfrozen soils to frozen soils?





### **Testing**

Phases	Purpose	Conditions	Samples
1	Prove the existence of an index similar to Pc	Same T Different γ <sub>d</sub>	4
		Same γ <sub>d</sub> Different T	4
2	Influence of creep on this index	$\gamma_d$ , T, preload	30
3	Comprehensive analysis Relationship: Creep vs. Pc	Orthogonal design: $\gamma_d$ , T, preload, creep time	10

#### 48 creep tests so far





K<sub>0</sub> Test

### Pseudo preconsolidation pressure <u>PPC</u> for frozen soils



In "ln(1+e)~lnP" coordinates, there is clearly such an index





## **PPC: Mechanical behavior of frozen soils**



PPC increases with the increase in γ<sub>d</sub>, then does not change obviously
 PPC increases linearly with the decrease in temperature
 Well reflects the bonding in frozen soils







>When preload is less than original PPC, PPC increases with time

>When preload is larger than original PPC, PPC decreases with time

We are not ready to get a relationship between PPC and temperature, creep time; but we successfully proved the existence of such an index.





# AN ELEMENT MODEL FOR CREEP OF FROZEN SOIL (Dr. Songhe Wang)







### Long-term load test

#### Field load test



#### Loading pile

#### Numerical model



#### Only creep of underlying permafrost was considered





### Numerical simulation

#### After implementation of this model,



A simple model for creep of frozen soil might provide a way in engineering analysis.





#### A VISCO-HYPOPLASTIC CONSTITUTIVE MODEL FOR FROZEN SOIL

(Dr. Guofang Xu, Prof. Wei Wu, Prof. Jilin Qi)

$$\dot{\sigma} = \dot{\sigma}_{s} + \dot{\sigma}_{d}$$

 $\sigma_{\rm s}$  is static stress,  $\sigma_{\rm d}$  is dynamic stress.

Static part

$$\dot{\boldsymbol{\sigma}}_{s} = c_{1}[\operatorname{tr}(\boldsymbol{\sigma}_{s}-\boldsymbol{c})]\dot{\boldsymbol{\varepsilon}} + c_{2}\frac{\operatorname{tr}[(\boldsymbol{\sigma}_{s}-\boldsymbol{c})\dot{\boldsymbol{\varepsilon}}]}{\operatorname{tr}(\boldsymbol{\sigma}_{s}-\boldsymbol{c})}(\boldsymbol{\sigma}_{s}-\boldsymbol{c}) + f_{\varepsilon}\cdot f_{cd}\cdot \left[c_{3}(\boldsymbol{\sigma}_{s}-\boldsymbol{c})^{2} + c_{4}(\boldsymbol{\sigma}_{s}-\boldsymbol{c})_{d}^{2}\right]\frac{\|\dot{\boldsymbol{\varepsilon}}\|}{\operatorname{tr}(\boldsymbol{\sigma}_{s}-\boldsymbol{c})}$$

in which c is the cohesion of frozen soil,  $f_{\varepsilon}$  is a scalar function of deformation,  $f_{cd}$  is a factor of creep damage.

$$f_{\varepsilon} = 2 - \exp(\alpha \cdot l + \beta)$$

 $\alpha$  and  $\beta$  are parameters, *l* is the accumulation of deformation.

$$l = \int_{t_0}^t \left\| \dot{\mathbf{\varepsilon}}(\tau) \right\| \mathrm{d}\,\tau$$





$$f_{\rm cd} = 1 + \gamma \cdot \int_{t_1}^{t_2} \langle \ddot{\epsilon}(\tau) \rangle d\tau$$

 $\gamma$  is a parameter,  $\langle \, \rangle \,$  is Macaulay brackets.

Dynamic part

$$\dot{\boldsymbol{\sigma}}_{d} = \eta_{1} \sqrt{\eta_{2}^{2} + \operatorname{tr}(\dot{\boldsymbol{\varepsilon}}^{2})} \ddot{\boldsymbol{\varepsilon}}$$

 $\eta_1$  and  $\eta_2$  are parameters,  $\mathbf{\bar{\epsilon}}$  is strain acceleration.

Complete constitutive model - Rate dependent

$$\dot{\boldsymbol{\sigma}} = c_1[\operatorname{tr}(\boldsymbol{\sigma}-\boldsymbol{c})]\dot{\boldsymbol{\varepsilon}} + c_2 \frac{\operatorname{tr}[(\boldsymbol{\sigma}-\boldsymbol{c})\dot{\boldsymbol{\varepsilon}}]}{\operatorname{tr}(\boldsymbol{\sigma}-\boldsymbol{c})}(\boldsymbol{\sigma}-\boldsymbol{c}) + f_{\varepsilon} \cdot f_{\mathsf{cd}} \cdot \left[c_3(\boldsymbol{\sigma}-\boldsymbol{c})^2 + c_4(\boldsymbol{\sigma}-\boldsymbol{c})_{\mathsf{d}}^2\right] \frac{\|\dot{\boldsymbol{\varepsilon}}\|}{\operatorname{tr}(\boldsymbol{\sigma}-\boldsymbol{c})} + \eta_1 \sqrt{\eta_2^2 + \operatorname{tr}(\dot{\boldsymbol{\varepsilon}}^2)} \ddot{\boldsymbol{\varepsilon}}$$





### CALIBRATION OF THE CONSTITUTIVE MODEL

#### Parameters in the static part

When the two linear and nonlinear terms in the static part of the model are abbreviated as  $L_1$ ,  $L_2$ ,  $N_1$  and  $N_2$ , this part can be rewritten as:

$$\dot{\boldsymbol{\sigma}}_{s} = c_{1} \mathbf{L}_{1}(\boldsymbol{\sigma}_{s}) : \dot{\boldsymbol{\varepsilon}} + c_{2} \mathbf{L}_{2}(\boldsymbol{\sigma}_{s}) : \dot{\boldsymbol{\varepsilon}} + c_{3} \mathbf{N}_{1}(\boldsymbol{\sigma}_{s}) \| \dot{\boldsymbol{\varepsilon}} \| + c_{4} \mathbf{N}_{2}(\boldsymbol{\sigma}_{s}) \| \dot{\boldsymbol{\varepsilon}} \|$$

In a conventional triaxial test, owing to  $\dot{\sigma}_2 = \dot{\sigma}_3 = 0$ , the above equation can be divided into two scalar equations as follows:

$$\dot{\sigma}_{1} = c_{1}L_{11}\dot{\varepsilon}_{1} + c_{2}L_{12}\dot{\varepsilon}_{3} + c_{3}N_{11}\sqrt{\dot{\varepsilon}_{1}^{2} + 2\dot{\varepsilon}_{3}^{2}} + c_{4}N_{12}\sqrt{\dot{\varepsilon}_{1}^{2} + 2\dot{\varepsilon}_{3}^{2}}$$
$$\dot{\sigma}_{3} = c_{1}L_{21}\dot{\varepsilon}_{1} + c_{2}L_{22}\dot{\varepsilon}_{3} + c_{3}N_{21}\sqrt{\dot{\varepsilon}_{1}^{2} + 2\dot{\varepsilon}_{3}^{2}} + c_{4}N_{22}\sqrt{\dot{\varepsilon}_{1}^{2} + 2\dot{\varepsilon}_{3}^{2}}$$





Owing to the radial stiffness  $E_{A3} = E_{B3} = 0$ , we have

$$\begin{split} E_{\rm A1} &= c_1 L_{11} + c_2 L_{12} v_{\rm A} + c_3 N_{11} \sqrt{1 + 2 v_{\rm A}^2} + c_4 N_{12} \sqrt{1 + 2 v_{\rm A}^2} \\ 0 &= c_1 L_{21} + c_2 L_{22} v_{\rm A} + c_3 N_{21} \sqrt{1 + 2 v_{\rm A}^2} + c_4 N_{22} \sqrt{1 + 2 v_{\rm A}^2} \\ E_{\rm B1} &= c_1 L_{11} + c_2 L_{12} v_{\rm B} + c_3 N_{11} \sqrt{1 + 2 v_{\rm B}^2} + c_4 N_{12} \sqrt{1 + 2 v_{\rm B}^2} \\ 0 &= c_1 L_{21} + c_2 L_{22} v_{\rm B} + c_3 N_{21} \sqrt{1 + 2 v_{\rm B}^2} + c_4 N_{22} \sqrt{1 + 2 v_{\rm B}^2} \\ A_{\rm A} = c_{\rm A} \frac{1}{v_{\rm A}} \frac{1}{v_{\rm B}} \frac{$$

The parameters  $c_i$  (i = 1, ..., 4) can be obtained by solving the above equation system with respect to the variables  $c_i$ .





### Parameters $\alpha$ and $\beta$

Parameters  $\alpha$  and  $\beta$  can be obtained from the following expressions:

$$\alpha = 1 - \left(T / T_{\rm ref}\right)^{n_{\rm l}}$$

$$\beta = -(T/T_{\rm ref})^{n_2}$$

in which  $T_{ref}$  is a reference temperature and can be regarded as -1°C (Zhu and Carbee, 1984).

Parameters  $\eta_1$  and  $\eta_2$  in the dynamical part can only be obtained by fitting the experimental data, as done by Hanes and Inman (1985).





# VERIFICATION OF THE CONSTITUTIVE MODEL Uniaxial compression tests at different loading rates

# $\varepsilon_{a}^{3500}$ $\varepsilon_{a}^{=1.1 \times 10^{-3} \text{ s}^{-1}}$ $\varepsilon_{a}^{=1.1 \times 10^{-4} \text{ s}^{-1}}$ $\varepsilon_{a}^{=1.1 \times 10^{-4} \text{ s}^{-1}}$ $\varepsilon_{a}^{=1.1 \times 10^{-5} \text{ s}^{-1}}$



Stress-strain relationship at different strain rates (Data from Zhu and Carbee, 1984)

#### Uniaxial creep tests at different stress levels







# **CONCLUDING REMARKS**

- General features in stress-strain-time curves for frozen are similar to that of unfrozen soils. Warm frozen soil is closer to unfrozen soils.
- A warm frozen soil is defined according to mechanical properties. Its creep was successfully observed in situ.
- Our own attempts: one is too simple, the other has too many parameters.
- No generally recognized constitutive models are found for creep of frozen soils. We have tried in different ways.

# Thank you for your attention!